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Current and emerging developments in subseasonal to decadal prediction

Article

Accepted Version

Merryfield, W. J., Baehr, J., Batte, L., Becker, E. J., Butler, A. H., Coelho, C. A., Danabasoglu, G., Dirmeyer, P. A., Doblas-Reyes, F. J., Domeisen, D. I., Ferranti, L., Ilynia, T., Kumar, A., Muller, W. A., Rixen, M., Robertson, A. W., Smith, D. M., Takaya, Y., Tuma, M., Vitart, F., White, C. J., Alvarez, M. S., Ardilouze, C., Attard, H., Baggett, C., Balmaseda, M. A., Beraki, A. F., Battacharjee, P. S., Bilbao, R., Marques De Andrade, F., DeFlorio, M. J., Diaz, L. B., Ehsan, M. A., Frangkoulidis, G., Grainger, S., Green, B. W., Hell, M. C., Infanti, J. M., Isensee, K., Kataoka, T., Kirtman, B. P., Klingaman, N. P. ORCID: https://orcid.org/0000-0002-2927-9303, Lee, J.-Y., Mayer, K., McKay, R., Mecking, J., Miller, D. E., Neddermann, N., Ng, C. H., Osso, A., Pankatz, K., Peatman, S., Pegion, K., Perwitz, J., Raclade-Coronel, G. C., Reintges, A., Renkl, C., Solaraju-Murali, B., Spring, A., Stan, C., Sun, Y. Q., Tozer, C. R., Vigaud, N., Woolnough, S. ORCID: https://orcid.org/0000-0003-0500-8514 and Yeager, S. (2020) Current and emerging developments in subseasonal to decadal prediction. Bulletin of the American Meteorological Society, 101 (6). E869-E896. ISSN 1520-0477 doi: 10.1175/BAMS-D-19-0037.1 Available at https://centaur.reading.ac.uk/88780/

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To link to this article DOI: http://dx.doi.org/10.1175/BAMS-D-19-0037.1

Publisher: American Meteorological Society

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Abstract

Weather and climate variations on subseasonal to decadal timescales can have enormous social, economic and environmental impacts, making skillful predictions on these timescales a valuable tool for decision makers. As such, there is a growing interest in the scientific, operational, and applications communities in developing forecasts to improve our foreknowledge of extreme events. On subseasonal to seasonal (S2S) timescales, these include high-impact meteorological events such as tropical cyclones, extratropical storms, floods, droughts, and heat and cold waves. On seasonal to decadal (S2D) timescales, while the focus broadly remains similar, (e.g., on precipitation, surface and upper ocean temperatures and their effects on the probabilities of high-impact meteorological events), understanding the roles of internal and externally-forced variability such as anthropogenic warming in forecasts also becomes important.

The S2S and S2D communities share common scientific and technical challenges. These include forecast initialization and ensemble generation; initialization shock and drift; understanding the onset of model systematic errors; bias correction, calibration, and forecast quality assessment; model resolution; atmosphere-ocean coupling; sources and expectations for predictability; and

linking research, operational forecasting, and end user needs. In September 2018 a coordinated pair of international conferences, framed by the above challenges, was organized jointly by the World Climate Research Programme (WCRP) and the World Weather Research Programme (WWRP). These conferences surveyed the state of S2S and S2D prediction, ongoing research, and future needs, providing an ideal basis for synthesizing current and emerging developments in these areas that promise to enhance future operational services. This article provides such a synthesis.

Capsule

Climate prediction on subseasonal to decadal time scales is a rapidly advancing field that is synthesizing improvements in climate process understanding and modeling to improve and expand operational services worldwide.

[Introductory text]

Beyond the tremendous progress in weather forecasting witnessed in recent decades (Bauer et al. 2015), predictive capabilities have expanded, increasingly seamlessly, to encompass climate on subseasonal to decadal time scales (Fig. 1 and Kirtman et al. 2013). These advances have been enabled by better observations, data assimilation schemes, and models originating both from the weather prediction and long term climate simulation communities, together with increased computational power supporting progressively higher resolution and larger ensembles that allow uncertainties to be better estimated and in some cases reduced.

International efforts under the auspices of the World Weather Research Programme (WWRP) and World Climate Research Programme (WCRP) have helped drive this progress through coordinated research to improve the accuracy and utilization of weather and climate predictions. Community research efforts under the WCRP led initially to climate predictions one to two seasons ahead becoming part of the World Meteorological Organization (WMO) operational infrastructure (Graham et al. 2011). More recently a joint WWRP and WCRP Subseasonal to Seasonal Prediction Project has started tackling the so-called weather-climate prediction desert from two weeks to a season (Robertson et al. 2018; Mariotti et al. 2018), aiming to underpin new WMO operations on those time scales (Vitart et al. 2017), and the NOAA-led SubX project has similar aims (Pegion et al. 2019). At longer ranges, WCRP-enabled research has quantified predictability from a year to a decade, and corresponding WMO operational infrastructure for annual-to-decadal climate prediction is now in place (World Meteorological Organization 2018; Kushnir et al. 2019).

As each of these efforts has progressed it has become increasingly apparent that common challenges exist across predictive time scales. These include understanding and adequately representing in models processes that give rise to predictability in the Earth system, consisting of the physical climate system—atmosphere, ocean, land and sea ice—together with associated biogeochemical cycling, especially of carbon (upper part of Fig. 1); capturing and communicating inherent uncertainties caused by the chaotic nature of weather and climate; correcting for and reducing imperfections in models that may systematically degrade forecast quality; and providing forecast information in a form that is applicable to decision making. At the same time, opportunities for usefully predicting elements of the Earth system beyond long-term means of standard meteorological variables , including land, ocean and sea ice properties and risks of weather extremes, have come into focus. The ultimate collective endeavor is to improve the prediction of the spatial–temporal continuum connecting weather to climate through a coordinated, seamless and integrated Earth system approach for the benefit of society.

In September 2018, international conferences[1] on subseasonal to seasonal prediction (S2S, encompassing forecast ranges from two weeks to a season) and seasonal to decadal prediction (S2D, encompassing ranges longer than a season, up to a decade) together with cross-cutting plenary

[1] The Second International Conference on Subseasonal to Seasonal Prediction (S2S) and Second International Conference on Seasonal to Decadal Prediction (S2D) were held 17-21 September 2018 at NCAR facilities in Boulder Colorado. These coordinated meetings involved 347 participants, including 92 early career scientists, from 38 countries, with a total of 368 oral and poster presentations. Further information including a complete list of contributions can be found at https://www.wcrp-climate.org/s2s-s2d-2018-home.

311 sessions were convened jointly by WWRP and WCRP. This represented a confluence of research
312 and operational climate prediction expertise and knowledge exchange across prediction time
313 scales that was unprecedented in scope. Selected outcomes, organized by themes
314 encompassing the challenges outlined above, are synthesized in this article.
315
316 **Mechanisms of predictability.**
317 *Subseasonal to Seasonal*
318 A major source of S2S predictability is the organization of tropical convection by the Madden
319 Julian Oscillation, or MJO (Woolnough, 2019), which is predicted skillfully by S2S project models
320 up to 3-4 weeks ahead (Vitart 2017). The MJO has worldwide impacts that depend on its
321 amplitude and phase, including modulation of tropical cyclone activity (Lee et al. 2018; Zhao et
322 al. 2019) and extratropical phenomena such as the East Asian summer monsoon (Li et al. 2018).
323 The associated tropical-extratropical teleconnections (Lin et al. 2019) impart S2S forecast skill
324 for many of these extratropical phenomena including Euro-Atlantic weather regimes, position
325 of the jet stream, atmospheric rivers (DeFlorio et al. 2019), and hail/tornado activity (Baggett et
326 al. 2018). However, good representations of the basic state both in the tropics and extratropics,
327 as well as tropical air-sea interactions and atmospheric convection (e.g., Yoo et al. 2015), are
328 necessary for these teleconnections to be correctly simulated by general circulation models
329 (Henderson et al. 2017).
330

S2S predictability also derives from the stratosphere through its relatively long time scales of variability[2] and lagged influences on the troposphere (Kidston et al. 2015). Interactions between the stratosphere and the troposphere from the tropics to the extratropics thus provide a promising source of S2S prediction skill (Butler et al. 2019). For example, in the winter Northern Hemisphere stratosphere the climatological westerly polar vortex exhibits extremes in variability, including sudden stratospheric warmings (SSWs) that are driven largely by Rossby waves from the troposphere. SSWs have lagged impacts on sea level pressure, surface temperature and precipitation, including pronounced tendencies for cold anomalies over northern Eurasia and warm anomalies over northeastern North America (e.g., Sigmond et al. 2013). Initializing forecasts during extreme stratospheric events provides increases in prediction skill of surface climate in such regions up to 3-6 weeks later (Domeisen et al. 2019c). However, the predictability of specific extreme stratospheric events is limited, ranging from a few days to about two weeks (Fig. 2) for different SSWs (Karpechko 2018; Taguchi 2018, Domeisen et al. 2019a), although models show evidence of under-confident forecasts in the stratosphere on S2S timescales (O'Reilly et al. 2019). Outstanding questions remain about the mechanisms of stratosphere-troposphere coupling processes, in particular on the causes, variability, and trends for the occurrence of SSW events (Ayarzaguena et al. 2018; Simpson et al. 2018) and why not all SSW events have similar downward effects (e.g., Garfinkel et al. 2013, Maycock & Hitchcock, 2015). In addition, further research is needed to assess the degree to which prediction models capture both the stratospheric variability and coupling processes.

[2] Including the quasi-biennial oscillation (QBO) of the tropical stratosphere, whose influences span a range of time scales and are addressed in the "Time scale interactions" subsection.

Among atmosphere-surface influences, land-atmosphere interactions have their greatest impact on subseasonal time scales in forecasts where land is initialized (Dirmeyer et al. 2018a), but also can contribute skill on weather prediction and multi-month time scales (Dirmeyer and Halder 2016, 2017). The most broadly impactful land attribute is soil moisture (Koster et al. 2004, 2016), but anomalies in soil temperature (Y. Zhang et al. 2019; Yang et al. 2019), snow cover (Jeong et al. 2012; Orsolini et al. 2013), and vegetation states (Williams et al. 2016) can all have significant impacts. A number of recent studies have focused on non-local impacts of land surface anomalies, showing for example that soil moisture anomalies can exert remote as well as local influences in boreal summer through driving of quasi-stationary Rossby waves and associated circulation anomalies (e.g., Teng et al. 2019; Wang et al. 2019). In addition, land surface and subsurface temperatures in spring may exert delayed downstream influences on precipitation (Xue et al. 2018), and evapotranspiration may remotely influence precipitation over land (Wei and Dirmeyer 2019).

Atmosphere-ocean interactions, fundamental for S2D predictability, can also be influential on S2S time scales. For example submonthly prediction skills for precipitation and temperature are enhanced over certain land areas including parts of Australia, the Maritime Continent and the contiguous United States when tropical sea surface temperature (SST) anomalies associated with El Niño Southern Oscillation (ENSO) are present (Hudson et al. 2011; Li and Robertson 2015; DelSole et al. 2017). Extratropical SST anomalies also can impart S2S skill through teleconnections, as shown for example by McKinnon et al. (2016) who identified a SST anomaly

pattern in the mid-latitude North Pacific that tends to precede heat waves and rainfall deficits in the eastern United States by up to 50 days.

Sea ice strongly influences surface fluxes and lower atmospheric temperatures particularly in the marginal ice zone, and provides a source of S2S predictability for polar and possibly midlatitude regions (Chevallier et al. 2019). This motivates the development of S2S forecasts for sea ice, which thus far have shown significant, albeit region-dependent skill for predicting intraseasonal Arctic sea ice variability (Liu et al. 2018, Zampieri et al. 2018).

Seasonal to decadal

A primary general source of S2D atmospheric predictability is remote influences from a variety of teleconnections (e.g., Yuan et al. 2018; Ruprich-Robert et al. 2018; Beverley et al. 2019. Teleconnections associated with anomalous atmospheric circulation patterns arise from changes to the Walker circulation usually driven by anomalous zonal SST gradients (Cai et al. 2019), and changes to the Hadley circulation usually driven by anomalous meridional SST gradients, especially interhemispheric differences (Kang et al. 2018). These influences impact tropical cyclones and rainfall, whereas anomalous upper level divergence due to tropical rainfall anomalies leads to Rossby waves that impact the extratropics (Scaife et al. 2017; O'Reilly et al. 2018). Besides giving rise to atmosphere-ocean interactions that alter the atmospheric circulation, SST anomalies can induce low-level temperature and moisture anomalies that are advected elsewhere by climatological winds (Dunstone et al. 2018).

S2D atmospheric predictability arising from teleconnections requires that SST anomalies be predictable. On seasonal timescales, tropical SST anomalies are dominated by ENSO (Yang et al. 2018), though there is some independent variability in the tropical Atlantic and Indian Oceans that also drives teleconnections (e.g., Nnamchi et al. 2015; Lim et al. 2016). The impacts of ENSO are sensitive to ENSO diversity (Capotondi et al. 2015), including the longitude at which maximum SST anomalies occur (Yeh et al. 2018; Patricola et al. 2018). ENSO SST anomalies are largely predictable out to a year particularly in winter and early spring (Barnston et al. 2017), whereas predictability may extend to two years for some La Niña events (Di Nezio et al. 2017), and to 1 ½ to two years for certain El Niño events (Luo et al. 2008).

Decadal SST variability occurs in both the Atlantic and Pacific oceans, often referred to as Atlantic Multidecadal Variability (AMV) and Pacific Decadal Variability (PDV), e.g. Kushnir et al. (2019). The causes of AMV are not fully understood, especially the relative roles of internal variability and external forcing from aerosols. However, AMV is modulated to some extent by the oceanic Atlantic Meridional Overturning Circulation (Yeager and Robson 2017), which together with the North Atlantic subpolar gyre is influenced by deep ocean density anomalies particularly in the Labrador Sea (Robson et al. 2016); these influences contribute to the especially high multi-year predictability in the North Atlantic (Buckley et al. 2019). AMV couples to the Hadley circulation, affecting hurricanes and Sahel rainfall as illustrated in Fig. 3 (Sheen et al. 2017), and can initiate atmospheric Rossby waves with remote influences including temperatures in parts of China (Monerie et al. 2018). AMV can influence PDV (Ruprich-Robert et al. 2017), and vice-versa. PDV may also be influenced by off-equatorial heat content

anomalies in the western Pacific Ocean (Meehl et al. 2016). Decadal variability of deep convection in the Southern Ocean influences temperatures in that region, potentially explaining recent increases in Antarctic sea ice (L. Zhang et al. 2019).

S2D atmospheric predictability also arises from longer time scale processes over land, mainly involving soil moisture (Chikamoto et al. 2017; Ardilouze et al. 2019) and vegetation (Weiss et al. 2014; Bellucci et al. 2015). These highlight the need for land surface initialization (Prodhomme et al. 2016a) and realistic vegetation models (Alessandri et al. 2017).

An additional source of S2D predictability is variations in radiative forcing, which provide significant skill on multi-year timescales (Smith et al. 2019). Much of this skill arises from changes in greenhouse gases, but anthropogenic aerosols may force decadal variations in AMV (Booth et al. 2012) and PDV (Smith et al. 2016; Takahashi and Watanabe 2016). Solar variability (Misios et al. 2019), and volcanic eruptions (Menegoz et al. 2018) including through their influence on ENSO (Khodri et al. 2017; Wang et al. 2018) and possibly AMV and the North Atlantic Oscillation (NAO; Swingedouw et al. 2017) affect climate on seasonal to decadal timescales and are potentially important sources of predictability. However, the relative roles of external radiative forcing and internal variability (W. Kim et al. 2018) continue to be explored.

Time scale interactions

438 The Quasi-biennial Oscillation (QBO) is a downward-propagating ~28-month oscillation of
439 easterly and westerly zonal jets in the tropical stratosphere, driven by upward equatorial waves
440 from the troposphere (e.g., Kim and Chun 2015). In addition to having high predictability and
441 some teleconnected influence on winter surface climate (e.g., Scaife et al. 2014a), the QBO
442 modulates the amplitude, persistence, and rate of propagation of the boreal wintertime MJO
443 (Fig. 4) through its impact on tropical convection via changes in static stability near the
444 tropopause (Yoo and Son 2016, Nishimoto and Yoden 2017). MJO amplitude is better predicted
445 at longer leads during the easterly phase of the QBO (Marshall et al. 2017), likely as a result of
446 longer persistence of the MJO rather than its greater initial amplitude (Lim et al. 2019).
447
448 The modulation of SSW probability of occurrence by tropical sources of variability, such as the
449 QBO, ENSO, or MJO, may extend probabilistic predictability of stratospheric variability to a few
450 months or longer if these relationships can be adequately captured by prediction models
451 (Garfinkel & Schwartz 2017; Garfinkel et al. 2018; Domeisen et al. 2019a,b).
452
453 There is increasing evidence of additional interactions between various sources of S2S and S2D
454 predictability across time scales. One example is that seasonal time scale variations in ENSO
455 modulate the MJO (Chen et al. 2016) and its impact on the NAO (Lee et al. 2019) with
456 consequent influences on weather over remote regions. Another is that ENSO teleconnection
457 to the extratropics has varied over multi-decadal time scales spanning the past 100+ years
458 (O'Reilly 2018), possibly modulating ability to predict the NAO (Weishiemer et al. 2019),

459 although sampling variability can also give rise to such long-term changes in teleconnections
460 (Yun and Timmermann 2018).
461
462 **Modelling issues.**
463 *Subseasonal to Seasonal*
464 Because S2S operational prediction is a relatively new enterprise, considerable efforts focusing
465 on fundamental aspects of forecast system design are occurring at operational centers
466 worldwide (Takaya, 2019). One major emphasis consists of methods to represent the
467 uncertainty in initial conditions (bred vector, singular vector, ensemble data assimilation) and
468 model physics (stochastic physics, Leutbecher et al. 2018). In addition, configurations of real-
469 time forecasts and hindcasts, including ensemble size, ensemble strategy (lagged ensemble
470 with different initial times or burst ensemble with a common initial time) and hindcast period,
471 impact forecast quality and ability to evaluate the performance of the hindcast. Identifying
472 suitable compromises and trade-offs in forecast system design is a challenge under practical
473 constraints for operational activities (costs, priorities, timeliness) and demands further
474 research.
475
476 From the modelling perspective, multiple operational centers are moving towards a unified, or
477 "seamless" coupled forecast system that can be applied across timescales from days to seasons
478 or longer. More S2S models are incorporating ocean and sea-ice components, and becoming
479 increasingly complex and complete in representing coupled processes in the Earth system. On
480 the other hand, poor representation of model physics, in particular clouds (Morcrette et al.

2018), results in model drifts and biases in surface land and ocean temperatures, which is a long-standing modeling issue that can degrade the skill of S2S predictions (Vitart and Balmaseda, 2017). Improvements in cloud parameterizations (Stan and Straus, 2019) and in representing the diurnal cycle of the atmospheric boundary layers are crucial for advancing S2S modeling. The Earth system modeling approach poses another challenge to initialize the ocean and sea ice components with high accuracy; for example there is a relatively large dispersion of initialized sea ice fields in current S2S models (Chevallier et al. 2017, Zampieri et al. 2018). Another important S2S modeling issue is predicting the MJO, owing to its importance as a source of subseasonal predictability (H. Kim et al. 2018). Multi-model evaluations have shown that S2S models have difficulties in representing MJO propagation across the Maritime Continent. Process-oriented diagnostics (Maloney et al. 2019) have identified a dry bias in the lower troposphere as one of the causes for the poor MJO propagation through weakening the horizontal moisture gradient over the Indian Ocean and western Pacific (Lim et al. 2018) and dampening the organization and propagation of the MJO. A recharge process whereby moisture builds up in the lower troposphere during the suppressed convection phase of the MJO, and that is key for MJO propagation around the Maritime Continent in boreal winter, is underrepresented in S2S models due to the dry bias (Kim 2017). Ocean coupling is another important process for the MJO (DeMott et al. 2015), and several studies have demonstrated that ocean coupling can improve MJO propagation and enhance predictive skill in models.

Poor vertical resolution, low model lid height, inadequate orographic and non-orographic gravity wave parameterizations, and biases in the tropospheric mean state (e.g., the location of

503 stationary Rossby waves) could limit the predictive skill from stratosphere-troposphere
504 coupling processes (Tripathi et al. 2015; Butler et al. 2016), but new generations of prediction
505 systems have rapidly improved in many of these areas. Future model development could
506 prioritize improved representation of orographic and non-orographic gravity wave drag and an
507 internally-generated QBO (Butchart et al. 2018). Better understanding of stratosphere-
508 troposphere coupling processes and the role of the stratosphere on surface skill could be
509 gained through case studies and stratospheric nudging experiments (Hansen et al. 2017).
510 Improved observations of the stratosphere (e.g., aerosols and chemistry) and other climate
511 components may improve S2S predictions. Finally, there is potential for modeling of
512 stratospheric ozone chemistry which provides surface temperature predictability on S2S time
513 scales due to its influence on high-latitude stratospheric circulation anomalies together with
514 their lagged surface impacts (Stone et al. 2019). Although that may currently be too resource-
515 intensive due to the many species and reactions that must be modeled, emerging machine-
516 learning techniques may provide pathways for incorporating chemistry-climate information into
517 S2S forecasts (Nowack et al. 2018).
518
519 *Seasonal to decadal*
520 Modeling issues for S2D prediction naturally overlap with those for S2S prediction. However,
521 the longer time scales of S2D prediction lead to a greater emphasis on representing slower
522 climate variations such as ENSO and AMV, and greater attention to reducing model biases in
523 the ocean that may take months to years to develop. Increased model resolution can reduce
524 model biases as illustrated in Fig. 5 (Jia et al. 2015; Müller et al. 2018), and improve skill

(Prodhomme et al. 2016b; Schuster et al. 2019; Infanti and Kirtman 2019), although the greater computational cost is not always justified (Scaife et al. 2019). More fundamental strategies involve analyzing/understanding model biases, before attempting to correct them a priori or a posteriori. Such analysis methods include comparing hindcasts with observations and multi-decadal historical or other simulations to distill causation for model errors, such as in the tropical Pacific (Shonk et al. 2018) or Atlantic (Voldoire et al. 2019). Similarly, errors in modeled variability or teleconnection patterns can be characterized by examining their evolution with lead time. Model biases can be corrected both through simple methods such as statistical bias correction and anomaly coupling (Toniazzo and Koseki, 2018), and more complex methods such as supermodeling, through which multiple models exchange information during a climate simulation (Shen et al. 2016).

Performance of S2D predictions is strongly tied to initialization of model components beyond the lower atmosphere. For example, stratospheric initial conditions are a source of seasonal winter NAO skill (e.g., O'Reilly et al. 2019; Nie at al. 2019) as illustrated in Fig. 6, and ocean initial conditions are crucial for skillfully predicting ENSO (Balmaseda and Anderson 2009), as well as decadal variability in the subpolar North Atlantic (Yeager and Robson, 2017; Borchert et al. 2018). However, initialization using full-field observational values can lead to initial shocks affecting skill (Kröger et al. 2018) and in such cases initialization combining observed anomalies with the model's own climatology can be beneficial until underlying model errors can be reduced (Volpi et al. 2017). Basic initialization strategies continue to be an active research area particularly for decadal prediction (Brune et al. 2018), and methods extending to forecast runs

such as the ensemble dispersion filter which replaces the ensemble members with the ensemble mean every three months (Kadlow et al. 2017) are also being explored. Comparisons that apply different initialization methods to the same model can yield valuable insights (Polkova et al. 2019); further issues specific to the initialization of the land, ocean, and sea ice components are considered in the next section.

Tackling these diverse and persistent modeling issues effectively will require sustained effort, as simple model-specific solutions may not cure the underlying problems, and ideally this should involve coordination between the S2S/S2D prediction, climate modelling, and data assimilation communities.

Initialization issues.

Atmosphere initialization

Accurate atmospheric model initialization is a basic requirement for numerical weather prediction because atmospheric initial conditions are the primary source of predictability on time scales less than a week or two (Fig. 1). It is enabled by sophisticated data assimilation systems that are the result of decades of advancement (Bauer et al. 2015). Subseasonal and seasonal prediction systems generally initialize their atmospheric components by such means, with the additional requirement that historical observations must be assimilated similarly to produce reanalyses that are used to initialize hindcasts. Because in situ and remotely sensed atmospheric observations are relatively dense there is generally good agreement between different reanalyses for the modern era implying relatively low uncertainty at heights below

about 10 hPa, although temporal inconsistencies can result from changes in observing systems (Long et al. 2017). Because atmospheric initial conditions contribute less to predicability on multi-annual time scales, some decadal prediction systems do not initialize the atmosphere (e.g., Yeager et al. 2018).

Land initialization

Climatically important land variables such as soil moisture and snow can be initialized by driving land surface models with observed atmospheric fields (e.g., Koster et al. 2009; Sospedra-Alfonso et al. 2016a) or, more directly, assimilation of land observations principally from satellites (Bilodeau et al. 2016; Muñoz-Sabater et al. 2019; Toure et al. 2018). Yet predictability from land surface states is being harvested only to the extent that land initial conditions and the relevant processes are represented realistically in forecast models (Koster et al. 2011; Ardilouze et al. 2017). Historically, land surface and atmospheric models are developed separately and their coupled behavior is not calibrated or validated (Dirmeyer et al. 2019), so that coupled processes are often not represented accurately (Dirmeyer et al. 2018b).

There are also observational limitations. In situ measurements of soil moisture are of varying quality and uneven distribution, and are not designed for real-time operational use (Dorigo et al. 2011). Satellite soil moisture monitoring (Entekhabi et al. 2010; Kerr et al. 2010), provides either very shallow or total column measurements including groundwater (Li et al. 2012), and is subject to uncertainties caused by vegetation, etc. (Al-Yaari et al. 2017). By contrast, soil moisture in forecast models is mainly a gross reservoir for the surface water balance, and its

591 variations do not represent all of the observed processes, particularly at sub-grid scales.
592 Therefore model soil moisture is only a crude representation of reality, although it still contains
593 useful information that can be largely consistent across different land models (Koster et al.
594 2009).
595
596 Climate forecasts can be improved by making high-quality land state observations an
597 operational priority for real-time reporting, and planning for long-term continuity in satellite
598 monitoring (Balsamo et al. 2018). This includes vegetation, especially as its interannual
599 variability and cycles of agricultural planting and harvest are not represented and can affect
600 surface fluxes and predictions (Alessandri et al. 2017). In addition, realistic snow initialization
601 can positively impact subseasonal predictions of surface temperatures (e.g., F. Li et al. 2019).
602 Along with coupled land-atmosphere model development (Santanello et al. 2018), such efforts
603 would facilitate improved predictions on weather to subseasonal time scales, as demonstrated
604 by numerous forecast model-based sensitivity studies such as that of Koster et al. (2011).
605
606 *Ocean and sea ice initialization*
607 The importance of initializing the oceans stems from their relatively long thermal and dynamical
608 time scales, through which they play an essential role in S2D climate predictability (Cassou et al.
609 2017). In addition, the oceans can influence S2S variability, for example through air-sea
610 interactions affecting the MJO (DeMott et al. 2015) and mesoscale eddy impacts on
611 atmospheric circulation (Saravanan and Chang 2019). Predicting future ocean evolution,
612 especially on S2D time scales, requires estimates of 3D ocean states for initialization. This in

613 turn requires a data assimilation method (usually in conjunction with a dynamical model) to
614 constrain ocean state estimates based on available observations. Similar considerations apply
615 to state estimates of sea ice. Comparisons of different ocean and sea ice state estimates as in
616 Fig. 7 can point to variables and regions for which they are most robust, as well as to where
617 uncertainties are relatively large (Balmaseda et al. 2015; Chevallier et al. 2017). Observing
618 system experiments in which certain observations are withheld have shown for example that
619 data from tropical ocean moorings positively impacts state estimates even when Argo float
620 data is also available (Fujii et al. 2015).
621
622 Recent enhancements in observing capabilities are enabling improvements in ocean and sea ice
623 state estimates, potentially leading to more accurate initial conditions and hence better
624 forecasts. For example, assimilation of satellite measurements of sea surface salinity (SSS) leads
625 to improvements in tropical Pacific ocean states and ENSO forecasts in experiments using an
626 intermediate-complexity coupled model (Hackert et al. 2019), whereas assimilation of satellite-
627 derived sea ice thickness (SIT) measurements has shown potential for improving sea ice
628 forecasts in operational seasonal forecasting systems (Chen et al. 2017; Blockley and Peterson,
629 2018). A major limitation is that these data sources have been available for less than a decade,
630 whereas considerably longer hindcast periods are required for forecast post-processing and skill
631 assessment, and temporal consistency of observational data used for initialization is required to
632 avoid artificial biases between hindcasts and forecasts. Forecasts thus continue to be initialized
633 typically without assimilation of SSS or SIT, from initial conditions that deviate appreciably from

634 available observations especially for SIT (Uotila et al. 2019). This motivates alternative
635 approaches for initializing SIT over multidecadal hindcast periods (Dirkson et al. 2017).
636
637 *Coupled data assimilation*
638 The atmosphere, land, ocean and sea ice components of climate prediction models have often
639 been initialized individually, without coupling. However, such an approach does not make
640 optimal use of observations, which may exert influences across the interfaces of the model
641 components. In addition, physical inconsistencies between the separately initialized
642 components may lead to rapid adjustments, or shocks. To overcome these limitations attention
643 has increasingly turned toward developing coupled data assimilation methods that treat
644 multiple components, such as atmosphere and ocean, simultaneously using observations from
645 each (Penny and Hamill 2017). Such methods are termed weakly or strongly coupled (Penny et
646 al. 2017). Weakly coupled methods apply assimilation independently to each model component
647 within the coupled model, so that the components may exchange information across their
648 interfaces. Such techniques have shown promise for reducing shocks (Mulholland et al. 2015),
649 and have begun to be applied operationally (e.g., Browne et al. 2019). Strongly coupled
650 methods apply assimilation to multiple model components in an integrated manner, so that
651 observations assimilated in one component can directly influence others. Such methods
652 remain experimental and thus far have been applied mainly in simplified models (e.g., Penny et
653 al. 2019).
654
655 **Ensemble predictions and forecast information.**

Subseasonal to Seasonal

In contrast to ensemble weather forecasts, a consolidated verification strategy for S2S predictions is not yet established, and developing such a framework that encompasses important forecast attributes such as accuracy, association, discrimination, reliability, and resolution has thus emerged as a priority (Coelho et al. 2018). (Accuracy measures error, or distance between forecast and observed values; association measures strength of the linear relationship between forecast and observation as through temporal or spatial correlations; discrimination measures by how much forecasts differ given different outcomes; reliability measures how well forecast probabilities correspond to observed frequencies of occurrence; resolution measures by how much outcomes differ given different forecast probabilities. Forecast quality encompasses all these attributes, whereas skill indicates quality relative to some benchmark such as persisted anomalies or climatological probabilities.) As for seasonal predictions, a purpose of S2S hindcasts is to provide a larger sample for more confident verification statistics than real time forecasts because they cover more years. However, because S2S hindcasts are initialized using re-analysis and most often have a smaller ensemble size, their verification generally underestimates real-time forecast quality. Operational centres are encouraged to compute and monitor verification statistics based both on hindcasts and real-time forecasts.

As has been demonstrated for seasonal prediction, S2S multi-model ensembles (MMEs) generally outperform individual models (Vigaud et al. 2017; Pegion et al. 2019). Currently, the

677 S2S and SubX MME projects are providing testbeds for research[3] as well as a foundation for
678 operational use (Vitart and Robertson 2019; Pegion et al. 2019). One focus for exploiting such
679 datasets is developing calibration procedures, post-processing steps that improve the
680 properties of probabilistic forecasts, to enable S2S ensemble forecasts to provide reliable
681 probabilities for particular conditions occurring or thresholds being exceeded, especially for
682 extreme events. The varied current choices among S2S project modelling systems for hindcast
683 and near real time initialization dates, hindcast period and ensemble size is, however, limiting
684 advances in developing multi-model calibration and combination procedures. In addition, the
685 value of these datasets for research would be enhanced if more comprehensive stratospheric
686 data were to be available across models.
687
688 S2S ensemble forecasts have shown promise in providing useful predictions and early warnings
689 for high impact climate and weather events including severe heat waves and cold spells, as well
690 as regional probabilities of the occurrence of tropical storms as illustrated in Fig. 8 (Vitart and
691 Robertson 2018). Examples include severe cold conditions over Europe associated with the
692 negative phase of the NAO, whose useful predictability into week 3 is enhanced by tropical–
693 extratropical teleconnections resulting from MJO activity (Ferranti et al. 2018), and atmospheric
694 rivers, plumes of intense water vapor transport that often trigger weather and hydrologic
695 extremes and are especially predictable at lead times of 3 to 5 weeks during certain MJO and
696 QBO phase combinations (Baggett et al. 2017). While modest overall skill at ranges longer than

[3] Hindcast and near real-time forecast data are available from S2S at www.s2sprediction.net and from SubX at http://iridl.ldeo.columbia.edu/SOURCES/.Models/.SubX/.

697 a week has been found for S2S predictions of springtime Sahelian heat waves including
698 measures of heat stress, such conditions following a strong El Nino were accurately forecast,
699 pointing to the tropical Pacific as a source of predictability for extremes in that region (Batté et
700 al. 2018).
701
702 A global precipitation hindcast quality assessment of the S2S prediction project models (Fig. 9)
703 was performed by de Andrade et al. (2019). Sub-seasonal prediction quality is modulated by
704 the MJO, QBO, ENSO in the tropics, changes in large-scale flow in the extra-tropics and
705 stratospheric tropical and extratropical variability (Butler et al. 2019). It is therefore important
706 to estimate the predictive skill of such events and identify their impacts on predictions of
707 weather and weather extremes. Evaluating the conditional prediction quality associated with
708 the key low frequency variability modes is instrumental for better understanding S2S
709 predictability mechanisms. For example, MJO predictive skill in the S2S MME ranges between
710 12 to 36 days and is affected both by the MJO amplitude and phase errors (Vitart 2017; Lim et
711 al. 2018; H. Kim et al. 2018). Communicating these variations in forecast quality, including if the
712 forecasts are no better than climatology, is extremely important as users with such knowledge
713 can better utilize and benefit from the forecast information. Furthermore, capitalizing on
714 “windows of opportunity” when skill is especially high increases the value of S2S forecasts
715 (Mariotti et al. 2020), and motivates their frequent initialization (ideally daily).
716
717 *Seasonal to decadal*

718 Limited forecast quality in current S2D ensemble prediction systems motivates research
719 initiatives that focus on extracting skillful and reliable information from the large amounts of
720 forecast and hindcast data available to potential users[4].
721
722 One emerging theme of such research is that S2D prediction systems sometimes underestimate
723 the predictable signal (Eade et al. 2014; Scaife and Smith 2018). As a result, very large ensembles
724 that effectively filter out unpredictable noise demonstrat higher skill in predicting phenomena
725 such as the winter NAO (Scaife et al. 2014b; Dunstone et al. 2016) and seasonal to multi-annual
726 regional precipitation variations (Dunstone et al. 2018; Yeager et al. 2018) than was previously
727 thought possible. While very large ensemble sizes hold value for isolating weak predictable
728 signals, much smaller ensemble sizes are sufficient for skillful prediction of tropical SST, for which
729 signal to noise ratios are much larger (Zhu et al. 2015). The causes of unrealistically low modeled
730 predictable signals (sometimes called the "signal to noise paradox") remain under investigation.
731 Two hypotheses stemming from hindcast experiments are that winter NAO skill is enhanced by
732 skillful prediction of a QBO teleconnection that is too weak in models (O'Reilly et al. 2019), and
733 that transient eddy feedbacks are too weak in models (Scaife at al. 2019). Others based on simple
734 models suggest that the NAO predictable signal is too weak because climate models switch

[4] Seasonal hindcast data from the WCRP Climate-system Historical Forecast Project (CHFP; Tompkins et al. 2017) are available at http://chfps.cima.fcen.uba.ar/access.php, and from the North American Multi-Model Ensemble (NMME, Kirtman et al. 2014) including real-time forecasts at https://iridl.ldeo.columbia.edu/SOURCES/.Models/.NMME/. Decadal hindcast data from the WCRP Coupled Model Intercomparison Project Phases 5 and 6 are available via https://esgf-node.llnl.gov/projects/cmip5/ and https://esgf-node.llnl.gov/projects/cmip6/.

735 between NAO regimes too rapidly (Strommen and Palmer 2019), or exhibit less persistent NAO
736 variability than is observed (Zhang and Kirtman 2019).
737
738 In the case of the winter NAO which is a key source of variability over the mid-latitude North
739 Atlantic and Europe, another approach to extract relevant information from over-dispersive
740 ensembles that leads to improved skill is to subsample ensemble members that are close to a
741 "first guess" statistical prediction of the NAO (Dobrynin et al. 2018); subsampling has shown
742 potential for improving European summer forecasts as well (Neddermann et al. 2019).
743
744 Estimating and realizing the predictability of key modes of variability is still a major challenge at
745 S2D time scales. ENSO is considered one of the most predictable phenomena on multi-seasonal
746 time scales, but longer-range skill has been viewed as limited. However, multi-year ensemble
747 predictions have shown evidence of skill in predicting long-lasting La Niña events that follow
748 warm events up to 24 months ahead (DiNezio et al. 2017; Luo et al. 2017). Challenges in the
749 initialization of such longer time scale predictions remain, as evidenced by multi-year predictions
750 in which skill for SST and precipitation over land improves with lead time in some areas,
751 suggesting that short-term adjustments following initialization may tend to degrade skill (Yeager
752 et al. 2018).
753
754 Calibration of ensemble forecasts is a necessary step to reduce the areas for which S2D forecasts
755 are unreliable and potentially misleading. Combinations of several forecasting systems such as
756 the North American Multi-Model Ensemble (NMME, Kirtman et al. 2014) are now routinely used

to increase ensemble reliability and improve forecast skill. Several recent efforts have explored weighted multi-model calibration methods to combine ensembles from different models in order to improve probabilistic seasonal forecasts for temperature and precipitation anomalies as well as forecast of extremes (Becker 2017). Calibration methods have also been developed for ensemble decadal hindcasts to adjust both the bias and ensemble spread with a parametric dependency on lead time and initialization time (Pasternack et al. 2018). Such methods are found to improve both the conditional bias and probabilistic skill of decadal hindcasts.

Climate forecasts for decision making.

Subseasonal to Seasonal

Many decisions are made on the S2S forecasting timescale, which sits between weather forecasts and S2D climate outlooks; therefore the continued development of S2S forecasts has the potential to benefit many sectors of society (Fig. 10). S2S forecasting is a rapidly maturing discipline, with emerging recognition for both the need and the potential use of forecasts on this timescale (White et al. 2017). While S2S forecasts are increasingly being used in government as well as a range of sectors including agriculture, energy, finance, health and water resource management – more engagement between S2S forecasters and end users is needed to increase the wider awareness and uptake of S2S forecasts.

Although scientific knowledge gaps, computational capacity, and gaps in observations and modeling currently limit the use of S2S forecasts to some degree, by increasingly placing decision makers at the forefront of S2S forecast development, an improved dialogue between

779 S2S forecasters, developers and end users will accelerate the awareness and application of S2S
780 forecasts to real-world decision-making.
781
782 To support the increased use of S2S forecasts for decision-making, the following
783 recommendations were identified for action following the Boulder conference:
784 ● A summary of existing stakeholder case studies is planned to be created to demonstrate past
785 and ongoing 'success stories', and support better engagement with end users and
786 stakeholders. As the S2S forecast needs and associated performance varies greatly between
787 different sectors and users, the wider community is increasingly working together on the co-
788 design and production of S2S predictions in order to better meet user needs. Several
789 applications of S2S forecasts are now being developed in different disciplines, such as the
790 EU-funded S2S4E project in the energy sector, a quasi-operational excess heat outlook
791 system in the health sector (Lowe et al. 2016), and S2S hydrologic prediction in the water
792 management sector. These efforts need to be catalogued and disseminated to guide further
793 user-driven decision-support products, and to support the continued development of S2S
794 forecast, verification metrics and related services.
795 ● Systematically assessing the relative skill (or lack thereof) of forecasting a series of historical
796 high-impact events, such as heat waves, extreme rainfall events, or wildfires, on the S2S
797 timescale would be a useful way to help demonstrate the potential of S2S forecasts to
798 decision-makers across multiple sectors. At present, such case studies are often ad-hoc and
799 typically not published in the wider literature; however, a collaborative effort that brings
800 together a set of demonstrable case studies, involving both forecasters and end users, would

fill this gap. For example, a series of 'tailored narratives', or 'storylines' (approaches that construct stories of plausible, non-probabilistic climatic futures that relate to a specific person or sector to counter perceived barriers; e.g., Hazeleger *et al.* 2015), may aid in the understanding of what S2S forecasts may deliver in the future.

- To support the co-design, uptake and use of S2S forecasts, S2Sapp.net is currently being established as a new network of researchers, modellers and practitioners – an 'open to all' global community with a shared aim of exploring and promoting cross-sectoral services and applications of this new generation of forecasts from across government, academia, and the private sector.

Seasonal to decadal

Research efforts are assessing the value of S2D forecast information for many applications, and initiatives such as the WMO's Global Seasonal Climate Update[5] and Annual to Decadal Climate Update (Kushnir et al. 2019) are making such information more widely available. However, consultation with decision makers is essential in order to tailor forecast information to the needs and expectations of users.

Fisheries management is one application for which S2D forecast information holds promise (Tommasi et al. 2017). This is due to reasonable skill for ocean prediction in regions of interest, coupled with strong influences of S2D climate variability on fish populations. Case studies

[5] https://public.wmo.int/en/our-mandate/climate/global-seasonal-climate-update

employing fisheries management decision frameworks have shown that SST forecast information can potentially increase fishery yields while reducing the risk of population collapse from combined effects of environmental factors and overfishing. However, significant challenges remain for fully realizing this potential. These include a need for improved initialization and reduced model errors in key ocean regions such as the US Northeast continental shelf, dynamical downscaling in cases where important spatial scales are not resolved by global models, and sufficiently accurate observational data for hindcast verification on these scales. In addition, incorporating biogeochemistry and marine ecosystem components into prediction systems will expand their potential capabilities, while posing additional verification challenges.

Another current focus of application-oriented research is water management. Global climate prediction models have been shown to have skill in predicting the next winter season's snowpack throughout much of the western US, where spring snowmelt is an essential water resource (Kapnick et al. 2018; Sospedra-Alfonso et al. 2016b). Because temperature influences snowmelt and runoff efficiency, skill in seasonal temperature forecasts can provide improved skill for seasonal water supply forecasts in this region (Lehner et al. 2017). Seasonal forecast skill has also been demonstrated for monsoon rainfall (e.g., Jain et al. 2019) and drought (Hao et al. 2018) with potential to inform water management decisions in many regions of the globe. Decadal forecasts potentially can meet planning horizon needs but currently are less familiar to water managers than seasonal forecasts and long-term climate projections. Efforts to apply decadal climate information for water management decisions have included assessing the role of decadal modes of variability, and using statistically downscaled decadal predictions as hydrological model inputs.

Developing information that is credible and compatible with existing decision frameworks is an important consideration (Towler et al. 2018).

Additional sectors for which S2D forecasts are being assessed for decision making include agriculture (Klemm and McPherson, 2017), energy (demand & wind power generation, Clark et al. 2017; Lledó et al. 2019), tropical cyclone (Bergman et al. 2019) and coastal flooding (Widlansky et al. 2017) preparedness, Arctic marine transportation (Stephenson and Pincus 2018), wildfire risk (Turco et al. 2019), and food security (Funk et al. 2019).

Initiatives to develop and deliver climate forecast information range in scale from international, regional and national (e.g., Marotzke et al. 2016), to individual users, all of which aim to provide forecast information having practical value for decision makers. In all cases, it is crucially important that uncertainties are adequately quantified and conveyed in order to avoid any false sense of certainty and to build trust in forecast information providers, although sometimes this requires overcoming a preference of users for deterministic information. Additional considerations are that expectations of users need to be conditioned to generally modest levels of skill, but that this information can nonetheless be advantageous when applied consistently in the long term. The likelihood that climate forecast information gets used increases when efforts are made to build relationships with potential users, and dialogs are opened to enable forecast products to be co-designed (Kolstad et al. 2019).

Cross-cutting issues in S2S and S2D prediction.

865 *Initialization shock and model error*

866 Model biases are an endemic modeling issue that is common across S2S and S2D prediction
867 time scales and influence all aspects of the prediction systems – complicating ingestion of
868 assimilated observations, degrading skill, and necessitating post-processing steps such as bias
869 correction and calibration for product development and delivery.

870

871 Model biases begin to develop on fast time scales and lead to drifts from the climate
872 represented by the initial conditions to that of a model's biased equilibrium state. It has been
873 extremely hard to understand the mechanisms behind these drifts, and further, pathways for
874 their diagnosis are not clear although some progress is being made (Sanchez-Gomez et al. 2015;
875 Shonk et al. 2018; Voldoire et al. 2019). Such difficulties arise due to non-linear interaction
876 between various physical processes that are parameterized, and because biases could be non-
877 local in their origin. Long time scales before models' equilibrium states are attained make
878 understanding the causes of drifts even harder. The Boulder meeting recognized that the
879 S2S/S2D prediction community needs to pay particular attention to developing pathways for
880 understanding the onset of model biases and put together mechanisms (such as summer
881 schools) to train the next generation of scientists with interest and expertise in climate
882 modeling and model diagnostics.

883

884 Initialization shocks that arise from imbalances in initial states with respect to the formulation
885 of the model and can be caused by limitations of observations and data assimilation as well as
886 model biases were also recognized as a major issue, particularly in the context of decadal

887 predictions. Initialization shocks result in the degradation of initial information that may be the
888 primary source of predictability for the subsequent forecast. Even after considerable research
889 and investment in decadal predictions it is still not clear what may be best approaches, such as
890 between full field vs. anomaly initialization, to retain predictive information in the initial state
891 while minimizing the influence of initial shocks on the subsequent forecast. The continuing
892 prominence of model drift and initial shocks as important issues reinforces a long held
893 sentiment that these are outstanding problems that need to be studied more systematically if
894 progress in translating inherent predictability into prediction skill is to be made.
895
896 *S2S and S2D research interactions*
897 The examples of interaction among modes of variability across S2S and S2D time scales noted
898 earlier emphasize the fact that continued interaction and communication across the S2S and
899 S2D research communities will be important to make progress. Furthering our understanding of
900 time-scale interactions will require investments in process level understanding of these
901 phenomena and will not only benefit our understanding about their lower-frequency variations
902 but will also contribute to improved process level diagnostics of model simulations. Better
903 understanding of time-scale interactions is likely to require the use of a hierarchy of models,
904 such as simple linear models to investigate the characteristics of tropical-extratropical
905 interactions (including their influence on storm tracks), to diagnose possible causes for errors in
906 their representation in complex GCMs (Dias et al. 2019).
907

Another aspect of research interactions across time scales is quantifying the fidelity of models in S2S and S2D prediction as well as projections of climate on longer time scales based on their simulation and prediction of shorter time-scale phenomena. The advantage of such an approach is that much larger samples for predictions of shorter time-scale phenomena are available, and an assessment of the reliability of such predictions can be used to build confidence in prediction on longer time-scales. Theoretical basis for extrapolating the reliability of forecasts across different time scales may also require the use of a hierarchy of models (Weisheimer and Palmer 2014; Christensen and Berner 2019).

Research and operations

Post-processing to improve forecast quality is an important area of research that bears directly on operational activities. Post-processing is necessary because biases in forecasts can be as large as the predicted signal, and therefore require the use of bias correction and calibration techniques to adjust real-time predictions before their delivery to the users. These requirements are shared across sub-seasonal to decadal prediction time-scales, however because of different levels of experience (seasonal predictions having the longest history) the opportunity for cross-community interactions was recognized. Some aspects for post-processing are specific to time-scale, for example, bias correction for decadal predictions requires methods to account for the non-stationarity of climate, and research needs to develop such methods were stressed.

Necessity for post-processing requires an extensive set of hindcasts to accompany real-time predictions. Because of limited resources, decisions about hindcast period, ensemble size and forecast start dates are not straightforward and call for further guidance from the research community. Such questions about the operational infrastructure for long-range prediction systems, including ensemble generation techniques and recommendations for harmonizing hindcast and real time forecast production, provide an opportunity to link operational and research communities that was highlighted during the conference.

Product development and communicating forecasts to the user community is also a common thread across the S2S and S2D communities. Communication of probabilistic forecast information (including confidence in the forecast based on past verifications) to users for their decision making has been a challenge, and once again there is much to be gained from lessons learned from the experiences of different communities. Similar challenges and opportunities also exist in the context of product development that incorporate user needs based on an ongoing dialog from the very start of the process. In addition, users often wish to have information on smaller spatial scales than are represented in global climate models. For such applications either statistical or dynamical downscaling is possible and can be effective in reducing local climatological biases, although clear demonstrations that downscaling can improve the skill of climate predictions remain elusive (e.g., Manzanas et al. 2018).

In summary, research needs for further development of operational infrastructure, product generation and communication of probabilistic forecasts were themes often repeated during the conference.

Conclusions and the future of subseasonal to decadal prediction

This paper has outlined many commonalities in the prediction of weather and climate across time scales and Earth system components, and through the value cycle from basic research to operational delivery.

The Earth's weather and climate is inherently chaotic and challenges the best currently available modeling capabilities. There remains however untapped skill, and realizing this skill will require improvements on numerous fronts. These include fundamental understanding of fine-scale processes, leading toward their robust parameterization; accurately representing property exchanges across Earth system components through realistic coupling limiting systematic errors; sustained Earth observing systems and advanced data assimilation methods enabling balanced initial conditions that avoid shocks and mitigate model drifts; and innovative numerical and ensemble generation techniques to address model scalability issues. Additional important avenues toward improved services include development of probabilistic information for high impact weather and climate events including unprecedented extremes, and optimal post-processing and data fusion to add value to multi-model ensembles, among many others.

These challenges are broad but so are opportunities for steady progress, ranging from curiosity-driven science to the systematic model evaluation and improvement in a collaborative and open research/operational environment.

The joint WWRP-WCRP conferences in Boulder clearly demonstrated the benefit in bringing relevant stakeholders together to cross-fertilize their experience, knowledge, respective issues and working cultures, which will surely help frame a new and vibrant research portfolio, and inspire the next generation of science leaders to ensure that society has access to the best possible weather and climate prediction science.

ACKNOWLEDGEMENTS

The International Conferences on Subseasonal to Decadal Prediction on which this paper is based were sponsored by: US CLIVAR, NSF, UCAR, NCAR and its Climate and Global Dynamics Laboratory (CGD), NOAA's Climate Variability and Predictability (CVP) and Modeling, Analysis, Predictions and Projections (MAPP) Programs, Copernicus Climate Change Service, IPSL, and WWRP/WCRP's Subseasonal-to-Seasonal (S2S) Prediction Project.

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SIDEBAR 1:

Hindcast and forecast quality assessment (**or, "the unexamined life is not worth living").**

Besides helping to inform decision making, the careful assessment of forecast quality is critical to guiding forecasting improvements, but has many and varied considerations. Simply answering the question "is this forecast better than that one?" is complicated, as the appropriate skill metric or means for comparison is not always obvious. Some recent studies have focused on newer statistical methods for comparing one forecast to another. One relatively simple approach is the random walk test (DelSole and Tippett 2016), illustrated in Fig. SB1. This method is applicable to a wide range of measures such as a score based on the earth mover's distance metric (Düsterhus 2019), while also being robust and fair in its discrimination.

The utility of forecast assessment can be illustrated through two very different applications of seasonal forecasts: sea-ice and hurricanes. The assessment of seasonal sea ice forecasts is complicated by the low quality of sea-ice observations, but nevertheless reveals that initializing sea-ice thickness using observational data sets generally improves the prediction of Arctic sea-ice extent and edges (Blockley et al. 2018). Comparison of multi-annual forecasts of Atlantic hurricane activity obtained by direct tracking of storms in decadal hindcasts and through a hybrid approach combining predicted SSTs and observed statistical relations finds that each approach is skillful, especially hybrid forecasts based on a SST index for AMV (Caron et al. 2018).

A robust assessment of model performance should include the model's simulation of climate modes and teleconnection patterns such as ENSO, MJO and NAO, since they are major sources of predictability and errors representating their patterns or strength (e.g., Yang and DelSole 2012; Vitart 2017) can degrade skill in affected regions (Gleixner et al. 2017; Lu et al. 2017). In cases where modeled teleconnection patterns are imperfect, forecast skill may be improved by means of statistical methods that use model forecasts of relevant climate modes such as ENSO as predictors (e.g., Strazzo et al. 2019). It remains desirable, however, for models to improve so that their simulated teleconnection patterns are sufficiently realistic that such corrections are not needed.

SIDEBAR 2:

Frontiers in Earth system prediction.

New frontiers in S2D prediction have been enabled by Earth system models (ESMs, Flato 2011) that represent the carbon and other biogeochemical cycles in addition to the physical climate system. These frontiers include prediction of ocean and land carbon sinks and biogeochemistry and their important contribution to global carbon storage, as well as ecosystem services. The world's oceans are a fundamental regulator of global carbon storage and variability. The strength of ocean carbon uptake, together with uptake of carbon by the land, determines the fraction of anthropogenic emissions remaining in the atmosphere, and hence modulates present and future warming. Observed global mean ocean carbon uptake shows variability on decadal time scales that can be represented by ESMs in which physical climate variables are assimilated (H. Li et al. 2019).

ESM simulations indicate that internal variability of the ocean carbon uptake on decadal timescales is as large as the forced climate change trend (Li and Ilyina 2018), pointing to the potential importance and utility of decadal carbon cycle predictions. Decadal predictions from a number of ESMs are assessing the predictability of the ocean and land carbon sinks and other ocean tracers such as dissolved oxygen. These forecasts are part of the Decadal Climate Prediction Project (Boer et al. 2016) and international programs such as the World Climate Research Programme's Grand Challenge on Carbon Feedbacks (Ilyina and Friedlingstein 2016). Initial results based on individual models have demonstrated potential predictive skill, assessed through verification against the assimilating reconstructions that provide initial conditions, for ocean carbon uptake in certain regions such as the North Atlantic reaching up to 7 or more years (Li et al. 2016; Lovenduski et al. 2019), and skill in predicting actual variations estimated from observations (Fig. SB2) has been demonstrated (Li et al. 2019).

ESM-based studies also point to the drivers of this predictability. Air-sea CO_2 flux mainly varies due to pCO_2 changes in the ocean. While thermal influences on pCO_2 play a role in shorter term predictability, predictability beyond 3 years is maintained mainly by nonthermal influences of ocean circulation and biological modification of surface dissolved inorganic carbon and alkalinity (Li et al. 2019; Lovenduski et al. 2019).

Investigations in progress are finding potential for multi-annual prediction of additional biogeochemical fields such as net primary productivity and interior dissolved oxygen concentrations. In addition, potential predictability and skill for terrestrial carbon uptake have

1758 also begun to be assessed, with promising initial results (N. Lovenduski 2019, personal
1759 communication). These examples demonstrate that skillful multi-year prediction is likely
1760 achievable for biogeochemical and ecological Earth system components, and open prospects
1761 for the utilization of such information although significant challenges including the paucity of
1762 long term observational data for initialization and verification will need to be overcome.
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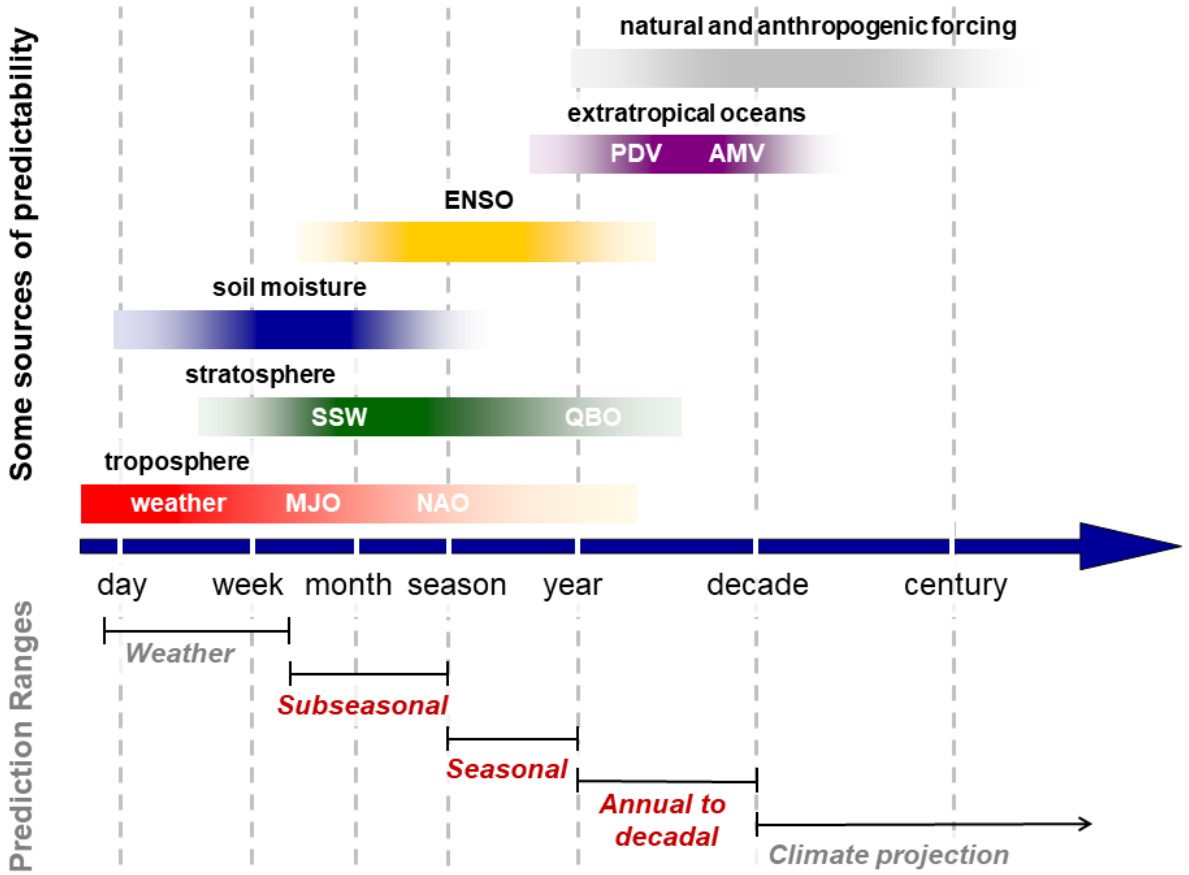

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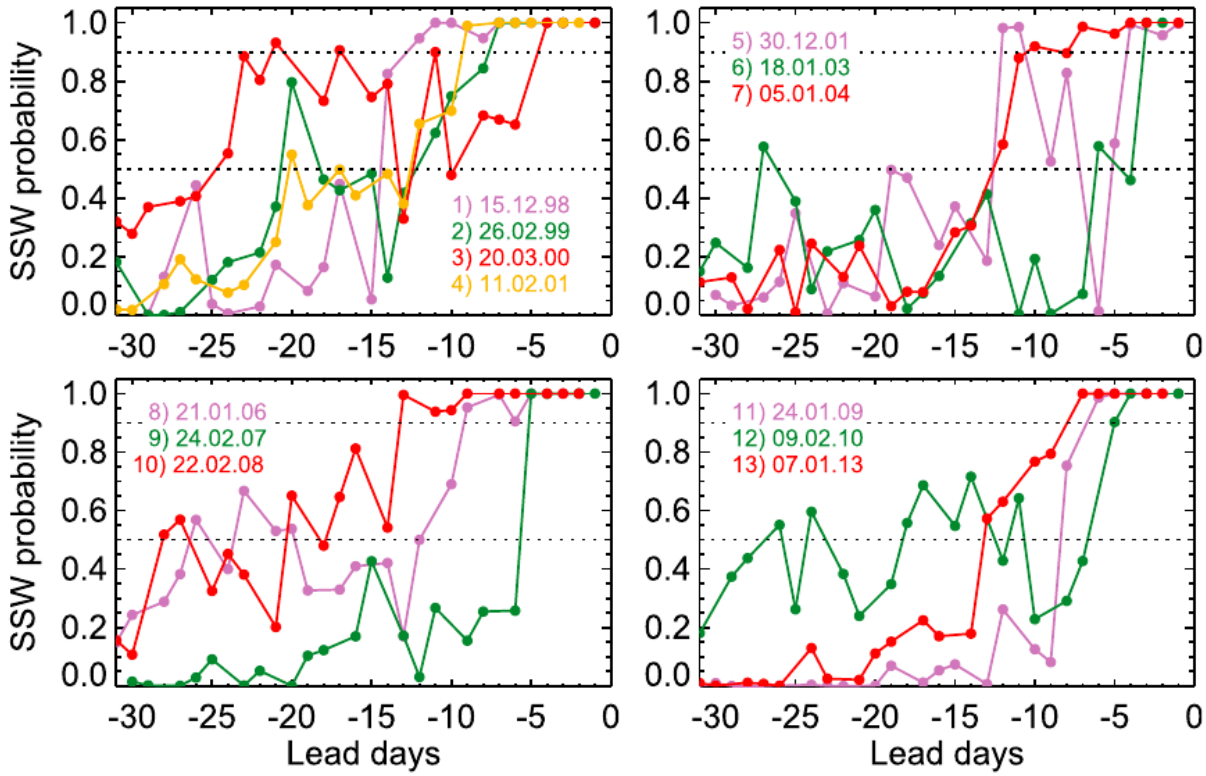
1771 **Fig. 1.** Schematic depiction of temporal ranges and sources of predictability for weather and climate
1772 prediction. The subseasonal range encompasses the S2S time scales, and the seasonal and annual-to-
1773 decadal ranges the S2D time scales, that are considered in this paper. Longer multi-decadal and
1774 centennial ranges derive predictability mainly from forcing scenarios rather than initial conditions, and
1775 are typically represented through climate projections originating from historical simulations begun in
1776 preindustrial times rather than predictions initialized from more recent observation-based climate
1777 states. Some important sources of predictability and approximate time scales over which they are most
1778 influential on surface climate are indicated in the upper portion of the figure; acronyms are defined and
1779 associated phenomena are discussed in the main text.



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1782 **Fig. 2.** Forecast probabilities of 13 SSW that occurred on the indicated dates as a function of lead time,

1783 based on ensemble hindcasts from the ECMWF monthly forecasting system. Most of the SSWs are

1784 predicted between 8 and 12 days lead time with a probability of 0.5–0.9, which is considerably larger

1785 than the average frequency of SSW occurrence. (From Karpechko 2018.)

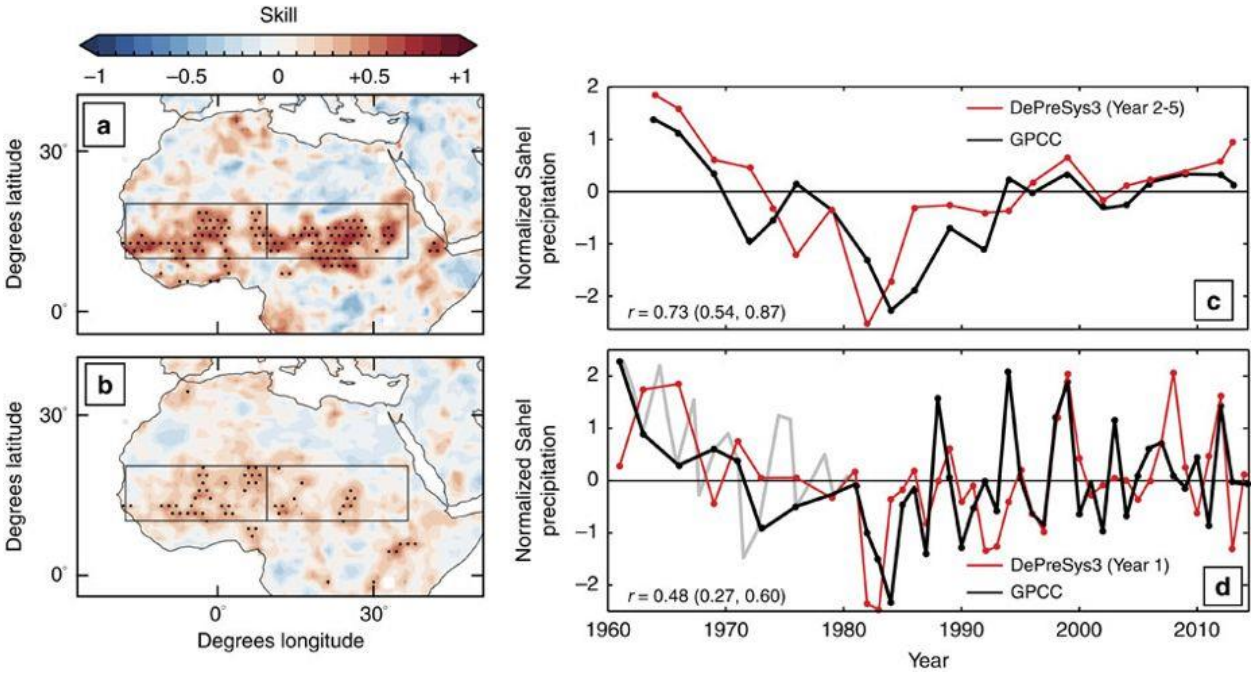


Fig. 3. Skill for predicting linearly detrended Sahel summer rainfall in years 2-5 (upper panels) and year 1 (lower panels) in DePreSys hindcasts. Panels (a)-(b) show spatial distributions of anomaly correlation coefficients with stippling indicating 95% significance. Panels (c)-(d) show time series of normalized predicted and GPCC observed rainfall anomalies in the Sahel region outlined by the boxes in the maps, with correlations and their 5–95% confidence intervals indicated. (From Sheen et al. 2017.)

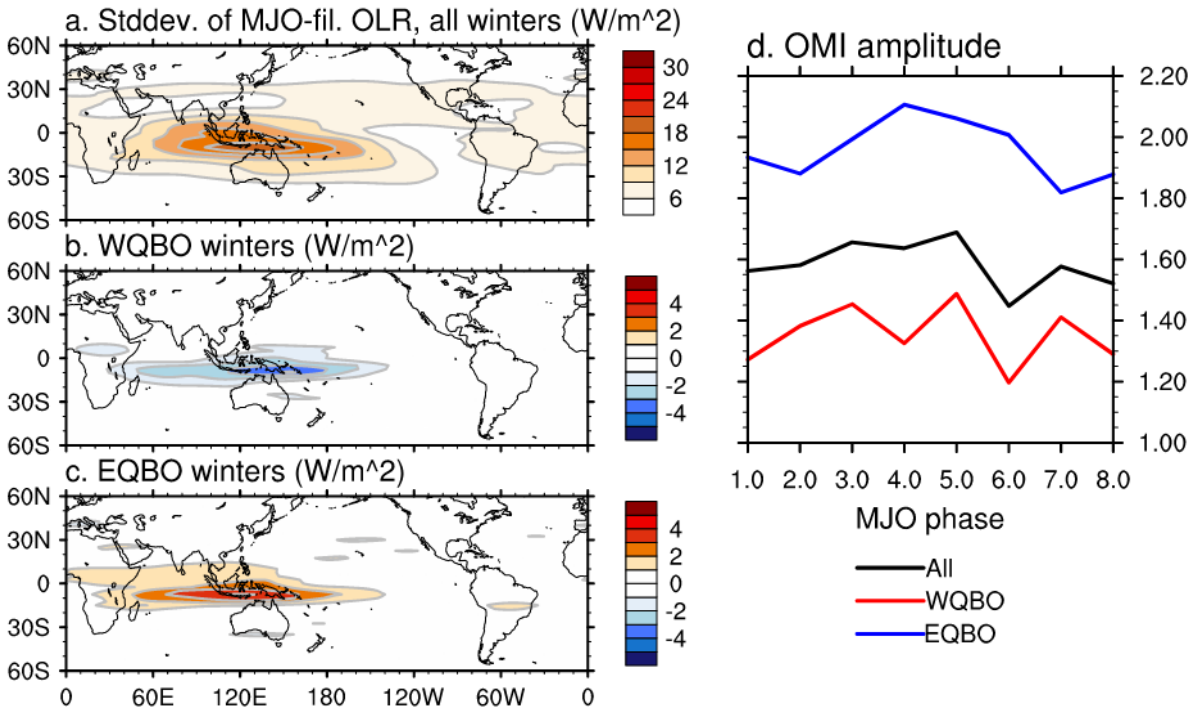


Fig. 4. Influence of QBO phase on MJO amplitude. (a) Standard deviation of wintertime outgoing longwave radiation (OLR), filtered to retain temporal and spatial scales characteristic of the MJO, for all winters in 1979 to 2012. Differences from these values in winters characterized by QBO westerly (WQBO) and easterly (EQBO) phases are shown (b) and (c) respectively. (d) Amplitude of an OLR-based MJO index (OMI) as a function of MJO phase for WQBO, EQBO and all winters. (From Yoo and Son 2016.)

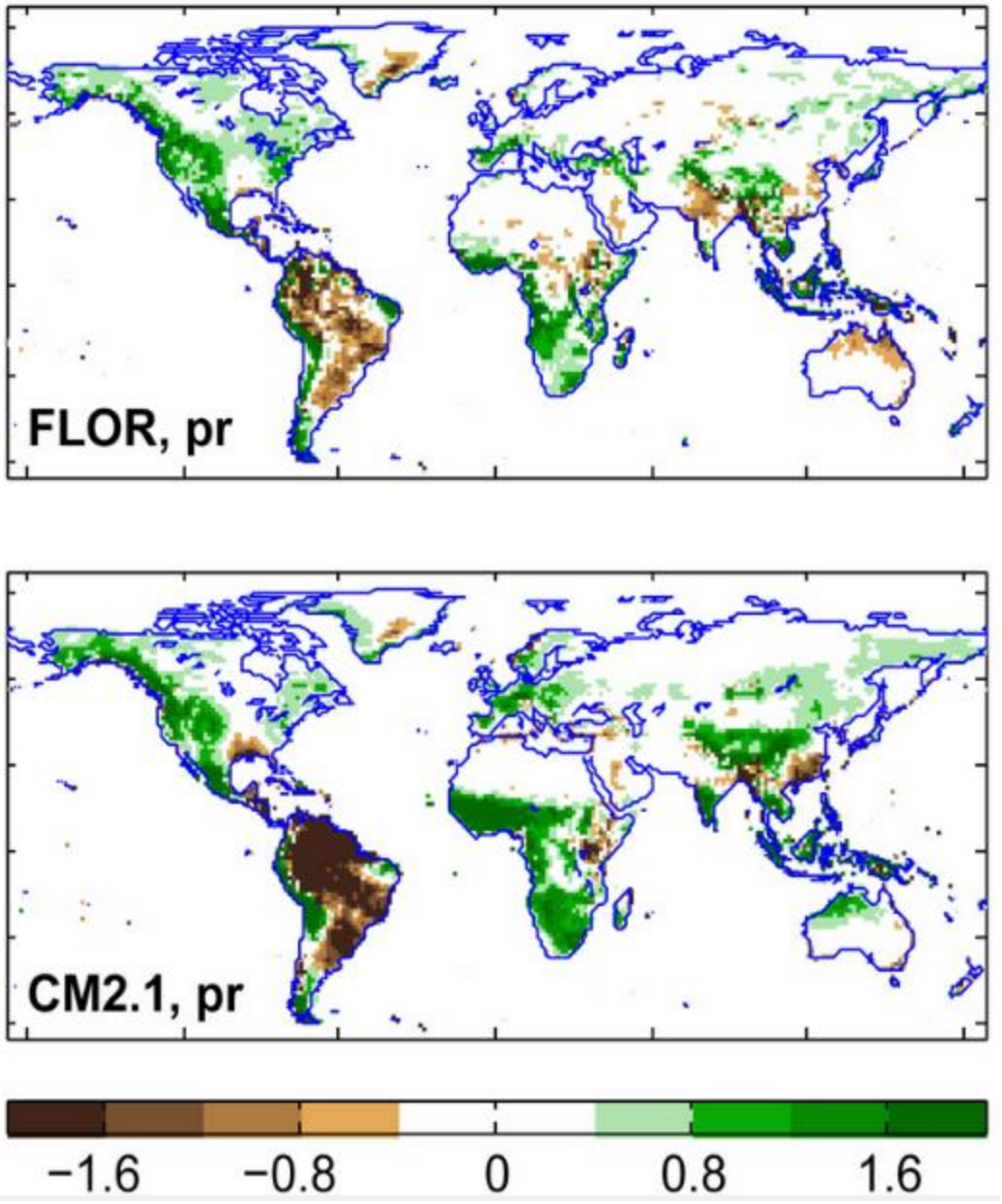
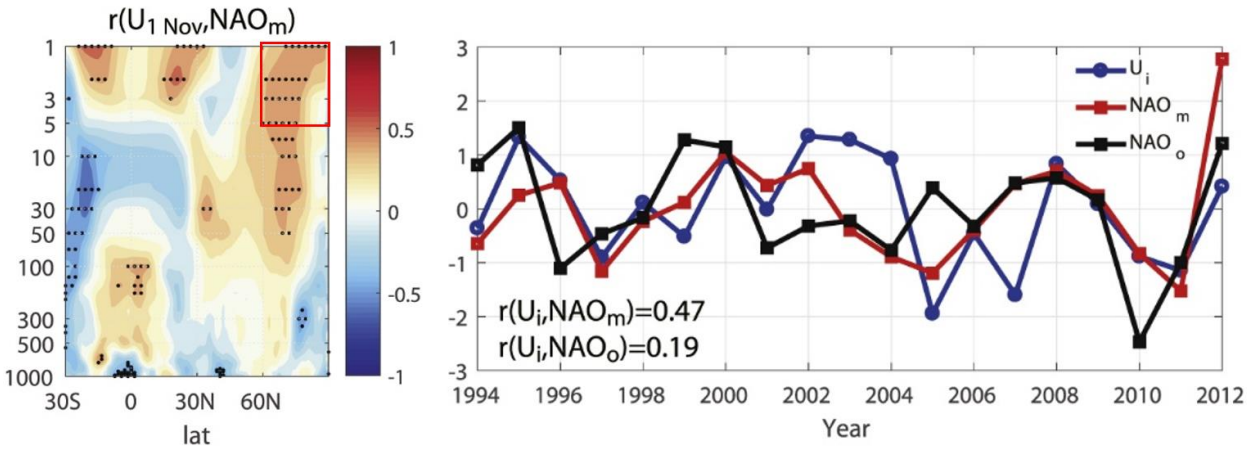


Fig. 5. Impact of resolution on precipitation biases in GFDL seasonal prediction models. Atmospheric resolution is approximately 50 km with 32 levels in FLOR (upper panel), and approximately 200 km with 24 levels in CM2.1 (lower panel), whereas ocean resolution is approximately 100 km in both models. Higher atmospheric resolution in FLOR reduces precipitation biases in numerous regions including much of the tropics. Annual mean biases over land in mm day^{-1} based on 1981-2010 CMAP observations are shown. (After Jia et al. 2015.)



1808

1809 **Fig. 6** Connection between stratospheric initial conditions and predicted winter NAO for UK Met Office

1810 GloSea5 predictions initialized 1 November 1995-2012. Left: correlation between initial zonal wind

1811 anomaly on 1 November and ensemble mean model-predicted surface NAO index (NAO_m) during DJF.

1812 Black dots represent values significant at $\alpha = 0.05$ confidence based on one-tailed test, and mean values

1813 within the red box define an index U_i. Right: Annual standardized U_i (blue), NAO_m (red) and observed

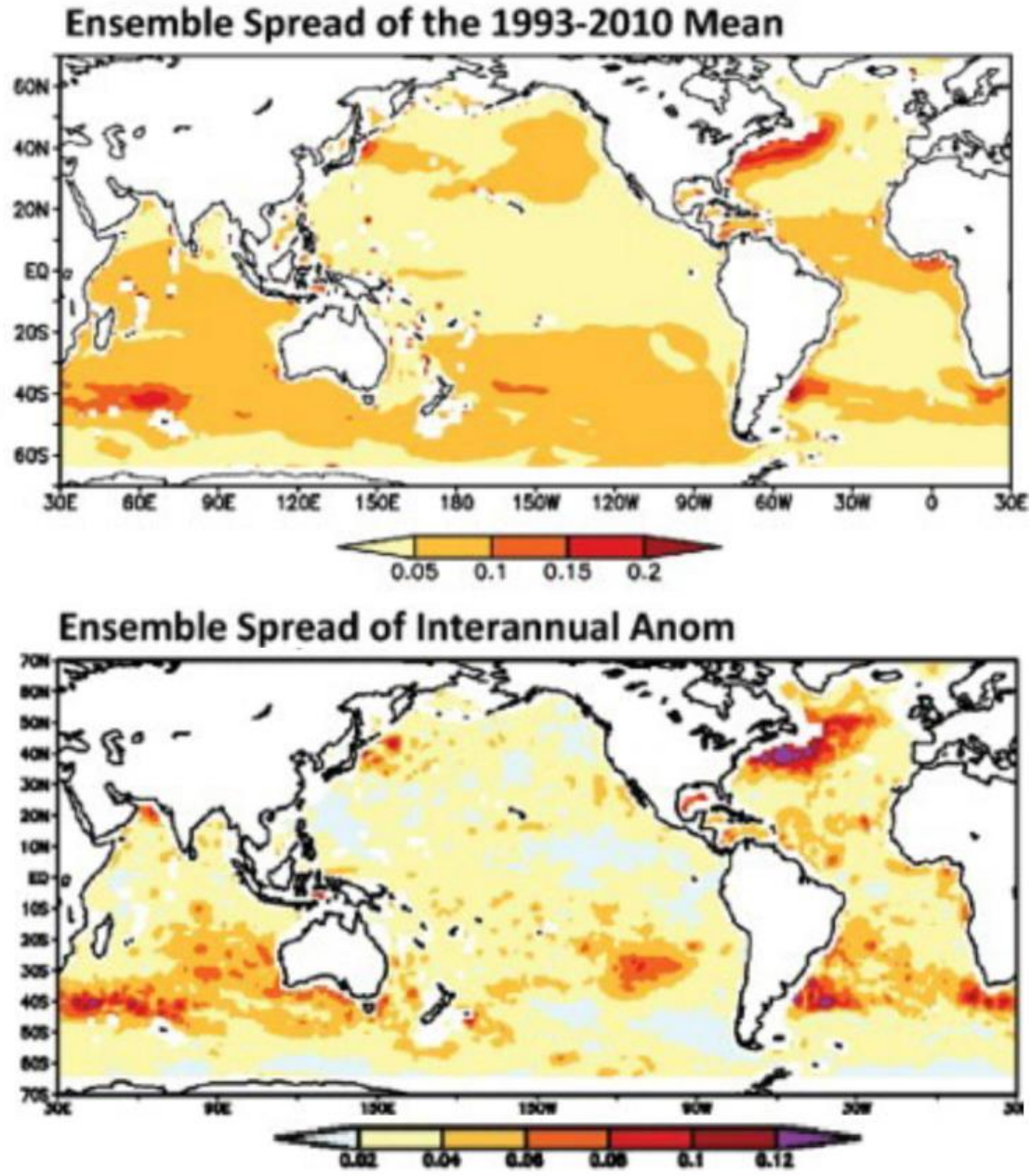
1814 surface NAO index, NAO_o (black). The correlation of U_i with NAO_m, indicated at lower left, is significant at

1815 $\alpha = 0.05$ confidence whereas the lower correlation of U_i with NAO_o is not unexpected based on signal to

1816 noise considerations and that there is only one realization of observations. The larger correlation of

1817 predicted and observed winter NAO values $r(NAO_m, NAO_o)=0.62$ suggests that additional sources of

1818 predictability exist. (After Nie et al. 2019.)



1819
1820 **Fig. 7.** Consistency across an ensemble of ocean state estimates of depth-averaged salinity over 0–
1821 700m, from the Ocean Reanalyses Intercomparison Project. Ensemble standard deviations in both the
1822 1993-2010 means (upper panel) and interannually varying monthly anomalies (lower panel) tend to be
1823 largest in eddy active regions such as the Gulf Stream, and less well-observed regions such as the
1824 Southern Ocean. These differences across state estimates are reflective of uncertainties in ocean initial
1825 conditions. (After Balmaseda et al. 2015.)

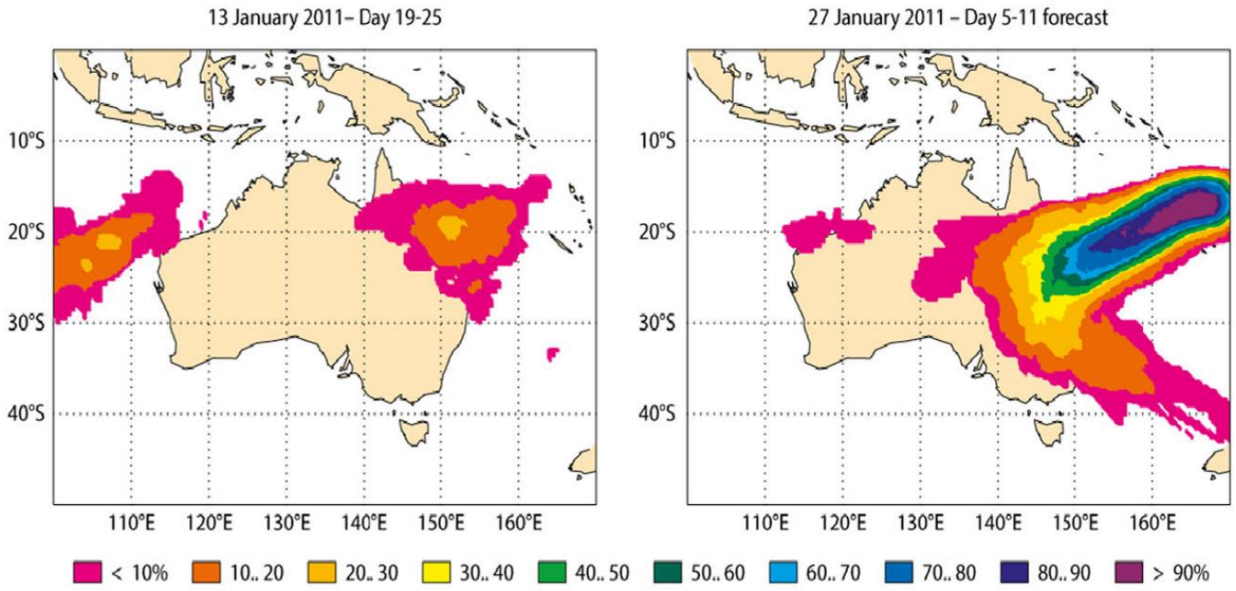
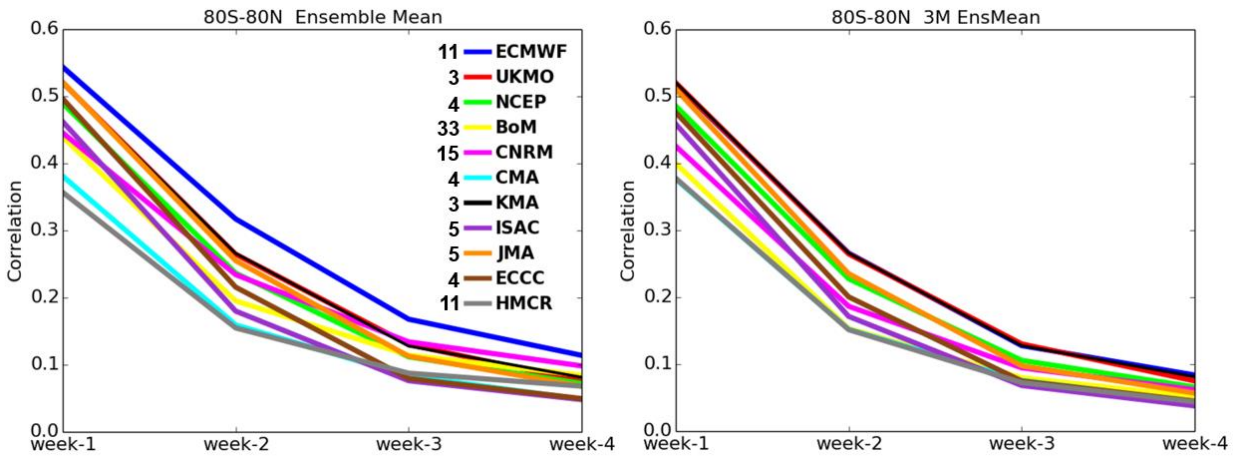


Fig. 8. Elevated probabilities of tropical cyclone occurrence during 31 January to 6 February 2011, based on ECMWF ensemble forecasts forecast starting 13 January with 18 day lead time (left), and 27 January with 4 day lead time (right). Destructive Cyclone Yasi made landfall in northeastern Australia on 3 February 2011 as a destructive category 5 storm. (Adapted from Vitart and Robertson 2018).



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1838 **Fig. 9.** Global averages of correlations between hindcast and observed precipitation anomalies over the
1839 80°S to 80°N latitudinal band for weeks 1-4 for S2S project models initialized from November to March,
1840 1999–2009. Left: Hindcast quality assessment based on ensemble means using the full ensemble size for
1841 each model, as indicated in the figure legend. Right: Hindcast quality assessment based on ensemble
1842 means using three ensemble members for each model. The reduced spread of the lines shown in the
1843 right panel when ensemble sizes are identical compared to the spread of the lines shown in the left
1844 panel demonstrates the value of the use of larger ensembles for subseasonal precipitation forecasting.
1845 (Adapted from de Andrade et al. 2019.)

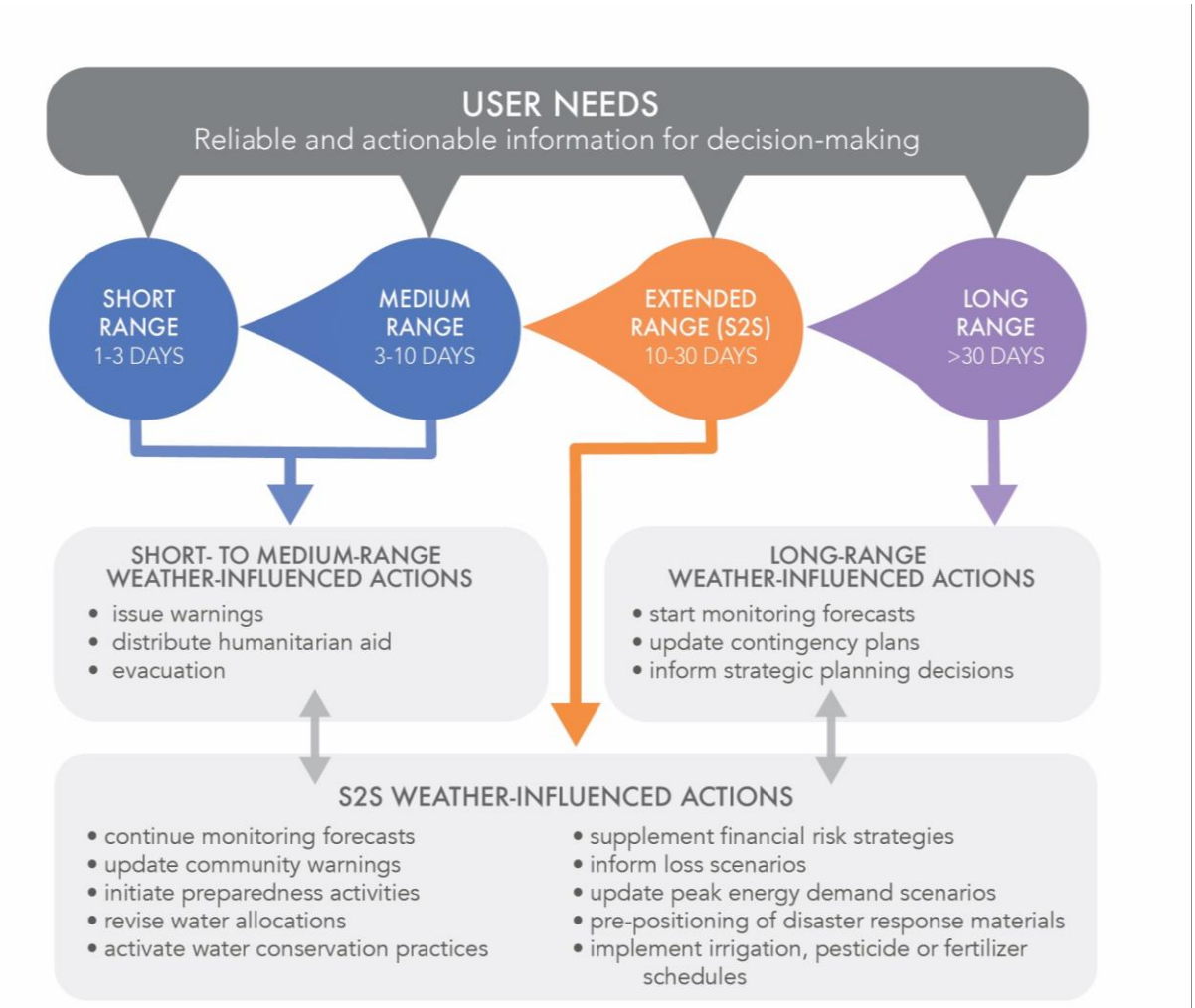

Fig. 10. Schematic illustration of relationships between a S2S forecast range of 10-30 days and other prediction timescales, including examples of actionable information that can enable decision making by various sectors. Indicated actions are examples that are not exclusive to a particular forecast range. (After White et al. 2017.)

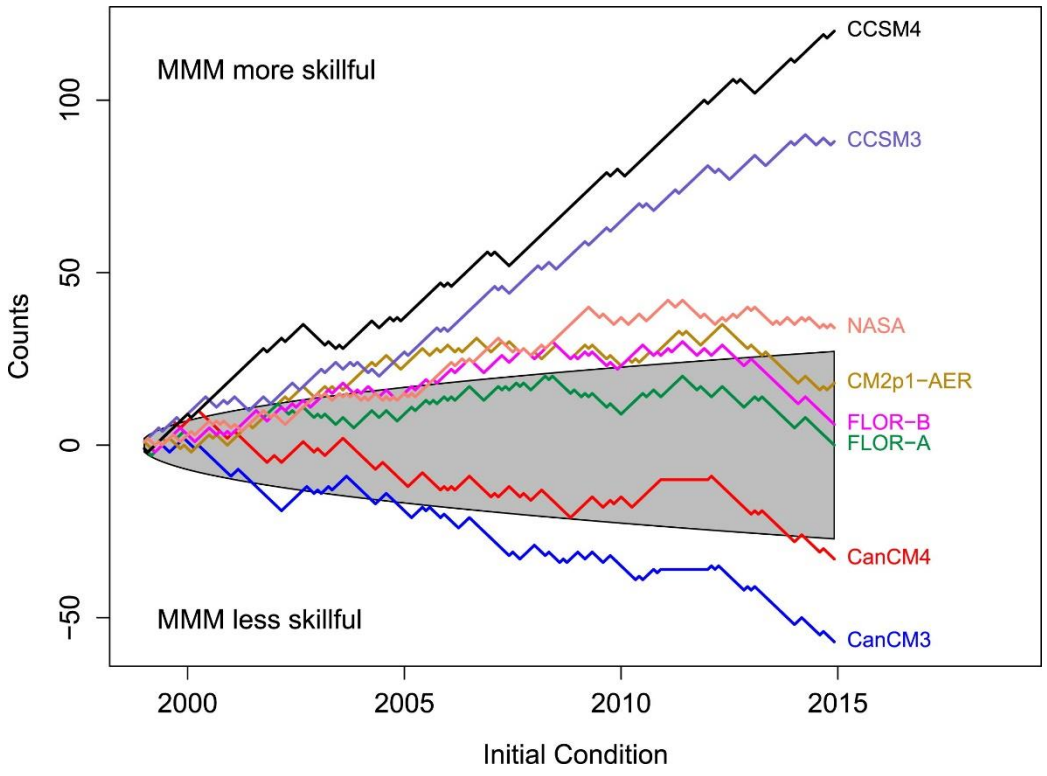

Fig. SB1. Random walk test comparing monthly mean forecasts of the Niño 3.4 index for equatorial Pacific SST at 2.5-month lead, between the multi-model mean (MMM) and individual models in the NMME. Counts (vertical axis) increase by 1 when MMM squared error is smaller than that an individual model (MMM more accurate) and decrease by 1 otherwise (individual model more accurate), and are accumulated forward for all initial months and years (horizontal axis). Accumulated counts above or below the shaded region indicate skill differences according to the squared error metric that are significant with >95% confidence (MMM more skillful above the shaded region and individual model more skillful below). Niño 3.4 anomalies are relative to 1982–98 climatological values, and span each month in 1999-2015. (From DelSole and Tippett 2016.)

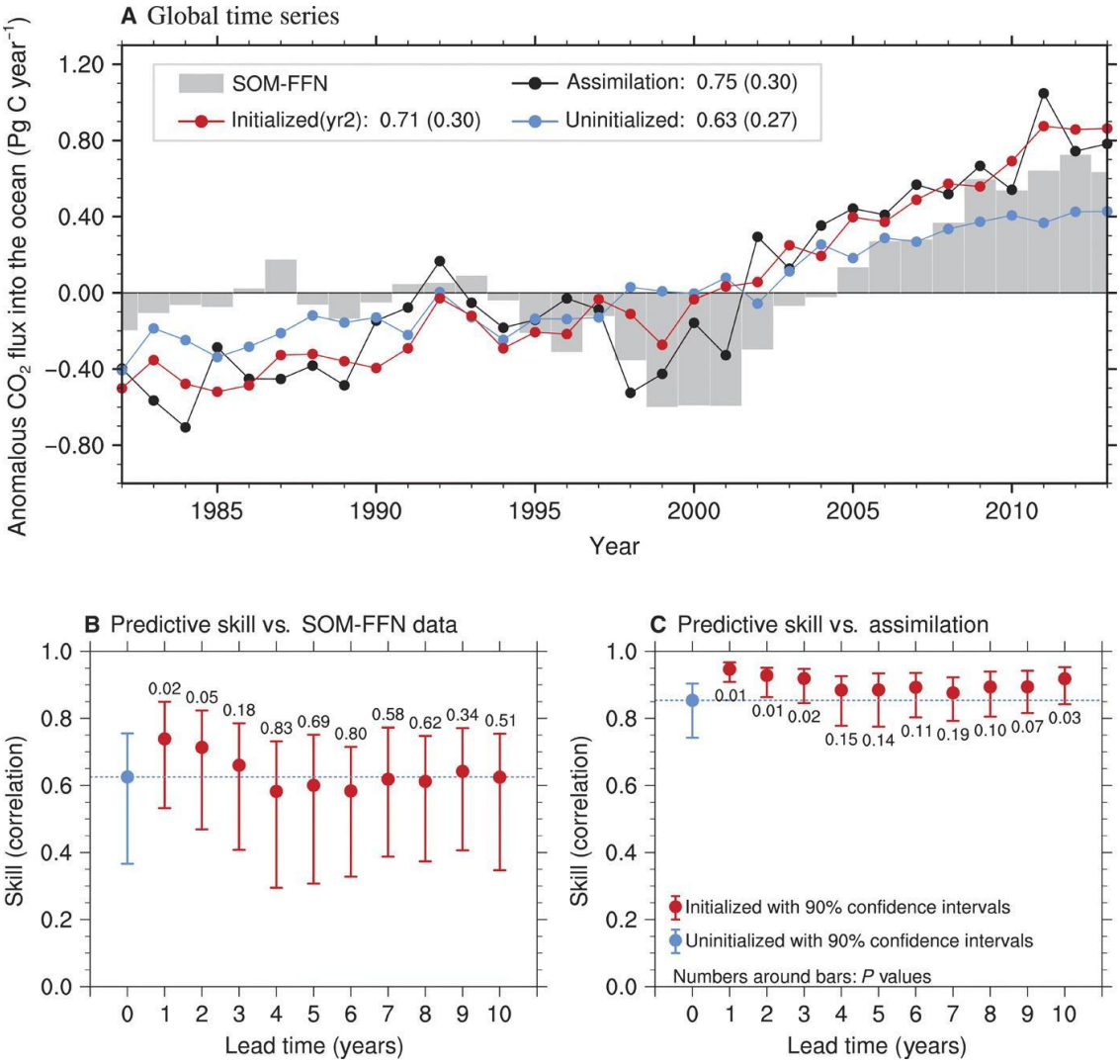

Fig. SB2. Temporal evolution and predictive skill of global CO_2 flux into the ocean from the MPI-ESM-HR decadal prediction system. (A) Annual values of anomalous CO_2 flux into the ocean from data–based estimates (SOM-FFN; gray) and MPI-ESM uninitialized simulations (blue), year 2 of initialized decadal predictions (red) and data-constrained assimilation run (black). Anomaly correlations and root-mean-square errors (in parentheses) verifying against SOM-FFN data are indicated. (B) Anomaly correlation skill for global CO_2 flux into the ocean, verifying against SOM-FFN. The blue dot and dashed line show the uninitialized skill for which lead time is not relevant, and the red dots initialized skill for different forecast years, with 90% confidence intervals and P values based on a bootstrap approach indicated. (C) Like (B), but verifying against the assimilation run. (After Li et al. 2019.)